

# Photoelectron balance in the dayside Martian upper atmosphere

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**Abstract:** Photoelectrons are produced by solar Extreme Ultraviolet radiation and contribute significantly to the local ionization and heat balances in planetary upper atmospheres. When the effect of transport is negligible, the photoelectron energy distribution is controlled by a balance between local production and loss, a condition usually referred to as local energy degradation. In this study, we examine such a condition for photoelectrons near Mars, with the aid of a multi-instrument Mars Atmosphere and Volatile Evolution data set gathered over the inbound portions of a representative dayside MAVEN orbit. Various photoelectron production and loss processes considered here include primary and secondary ionization, inelastic collisions with atmospheric neutrals associated with both excitation and ionization, as well as Coulomb collisions with ionospheric thermal electrons. Our calculations indicate that photoelectron production occurs mainly via primary ionization and degradation from higher energy states during inelastic collisions; photoelectron loss appears to occur almost exclusively via degradation towards lower energy states via inelastic collisions above 10 eV, but the effect of Coulomb collisions becomes important at lower energies. Over the energy range of 30–55 eV (chosen to reduce the influence of the uncertainty in spacecraft charging), we find that the condition of local energy degradation is very well satisfied for dayside photoelectrons from 160 to 250 km. No evidence of photoelectron transport is present over this energy range.

**Keywords:** Mars; Photoelectron; MAVEN

**Citation:** Wu, X. S., Cui, J., Yu, J., Liu, L. J., and Zhou, Z. J. (2019). Photoelectron balance in the dayside Martian upper atmosphere. *Earth Planet. Phys.*, 3(5), 373–379. <http://doi.org/10.26464/epp2019038>

## 1. Introduction

Photoelectrons are an important population of non-thermal particles in the upper atmospheres of Solar System planets such as Earth (Doering et al., 1976; Lee et al., 1980a, b), Mars (Shutte et al., 1989; Frahm et al., 2006a, b; Sakai et al., 2015), Venus (Coates et al., 2008, 2015; Tsang et al., 2015), and Titan (Coates et al., 2007; Wellbrock et al., 2012). They are produced by solar Extreme Ultraviolet (EUV) and X-ray ionization of atmospheric neutrals, usually denoted as primary ionization, as well as by impact ionization from primary photoelectrons with energy above the local ionization potential, usually denoted as secondary ionization (Fox et al., 2008, and references therein). Photoelectrons are of extensive aeronomical interest because they represent a key link between solar EUV/X-ray energy and atmospheric thermal energy (e.g. Chen RH et al., 1978; Choi et al., 1998; Matta et al., 2014).

On Mars, the typical photoelectron energy distribution is characterized by sharp peaks at 22–27 eV due to CO<sub>2</sub> and O photoionization by the strong solar He II emission line at 30.4 nm (e.g. Frahm et al., 2006a, b). These peaks are superimposed on a smooth and continuous background up to an apparent knee at 60–70 eV asso-

ciated with the rapid drop in solar irradiance at wavelengths shorter than 17 nm (e.g. Sakai et al., 2015; Peterson et al., 2016). With increasing energy, another less prominent peak structure occurs near 500 eV associated with inner shell Auger electrons due to X-ray ionization (e.g. Mitchell et al., 2000). Up to now, *in situ* measurements of photoelectrons have been made on various spacecrafts, such as the Mars Global Surveyor (e.g. Mitchell et al., 2000; Trantham et al., 2011; Liemohn et al., 2012; Xu SS et al., 2016a), the Mars Express (e.g. Frahm et al., 2006a, b, 2010; Han X et al., 2014), and the Mars Atmosphere and Volatile Evolution (MAVEN) (e.g. Sakai et al., 2015; Peterson et al., 2016; Xu et al., 2016b, 2017a, b), among others.

Upon production via either primary or secondary ionization, photoelectrons in the Martian upper atmosphere are subject to a number of collision processes including inelastic collisions with atmospheric neutrals leading to excitation or ionization (e.g. Bhardwaj and Jain, 2009), and Coulomb collisions with ionospheric thermal electrons (e.g. Stamnes and Rees, 1983). Assuming steady state, and at altitudes where the effect of transport is negligible, photoelectron production and loss should be in balance for any given energy and direction, a situation frequently referred to as local energy degradation (e.g. Mantas and Hanson, 1979).

This study is devoted to a verification of local energy degradation of photoelectrons observed in the Martian upper atmosphere. For

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Received 04 JUN 2019; Accepted 30 JUN 2019.

Accepted article online 03 JUL 2019.

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such a purpose, information regarding a number of controlling factors is necessary, such as the solar EUV/X-ray flux, the photoelectron intensity, the neutral and electron densities, as well as the electron temperature. Simultaneous measurements of the above parameters became available with the arrival of MAVEN at Mars on 21 September 2014 (Jakosky et al., 2015).

This paper is organized as follows. We describe in Section 2 the formulism used to calculate various photoelectron production and loss rates. In Section 3 we present the main results from our detailed calculations based on a multi-instrument data set accumulated over a representative dayside MAVEN orbit. Finally, conclusions are summarized in Section 4.

## 2. Formulism

The various photoelectron production and loss processes considered in this study include primary and secondary ionization (Nicholson et al., 2009), inelastic collisions with atmospheric neutrals (Bhardwaj and Jain, 2009), and Coulomb collisions with ionospheric thermal electrons (Stamnes and Rees, 1983).

Following Lavvas et al. (2011), the primary production rate,  $P_e^{(ine)}$ , is obtained from

$$P_e^{(pri)}(z, E) = \frac{1}{4\pi} \sum_{k,j} F_{\infty}(\lambda) e^{-\tau(z, \theta, \lambda)} \sigma_{k,j}^{(pi)}(\lambda) N_k(z), \quad (1)$$

where  $z$  is the altitude,  $E$  is the photoelectron energy,  $\lambda$  is the wavelength,  $F_{\infty}$  is the solar flux at the top of the atmosphere,  $\sigma_{k,j}^{(pi)}$  is the photoionization cross section for neutral species,  $k$ , and ionization channel,  $j$ ,  $N_k$  is the neutral density, and  $\tau$  is the optical depth along the line-of-sight with  $\theta$  being the solar zenith angle (SZA). For a given species and a given channel, the energy of the released photoelectron is  $E = hc/\lambda - E_{k,j}^{(th)}$  where  $h$  is the Planck constant,  $c$  is the speed of light, and  $E_{k,j}^{(th)}$  is the respective ionization potential. Similarly, the secondary ionization rate,  $P_e^{(sec)}$ , is obtained with

$$P_e^{(sec)}(z, E) = \sum_{k,j} \Phi_e(z, E') \sigma_{k,j}^{(ei)}(E') N_k(z), \quad (2)$$

where  $\Phi_e$  is the mean differential photoelectron intensity,  $\sigma_{k,j}^{(ei)}$  is the electron impact ionization cross section, and the energy of the released secondary photoelectron,  $E$ , is connected to the energy of the impacting primary photoelectron,  $E'$ , via  $E = E' - E_{k,j}^{(th)}$ .

Collisions with atmospheric neutrals cause both production and loss of photoelectrons. The respective rates, denoted as  $P_e^{(ine)}$  and  $L_e^{(ine)}$ , are given by

$$P_e^{(ine)}(z, E) = \sum_{k,j} \int_E^{\infty} \Phi_e(z, E') R_{k,j}(E', E) N_k(z) dE', \quad (3)$$

and

$$L_e^{(ine)}(z, E) = \sum_{k,j} \Phi_e(z, E) \sigma_{k,j}^{(ine)}(E) N_k(z), \quad (4)$$

where  $R_{k,j}$  is the redistribution function and  $\sigma_{k,j}^{(ine)}$  is the inelastic collision cross section for electron impact. The former reflects the probability that a photoelectron with energy  $E'$  is degraded to energy  $E$  via various channels of collisional relaxation.

The photoelectron loss rate via Coulomb collisions,  $L_e^{(cmb)}$ , can be calculated from

$$L_e^{(cmb)}(z, E) = -N_e(z) \frac{\partial [S(z, E) \Phi_e(z, E)]}{\partial E} \quad (5)$$

with  $N_e$  being the thermal electron density and  $S$  being the stopping cross section formulated as

$$S(z, E) = \frac{3.37 \times 10^{-12}}{E^{0.94} [N_e(z)]^{0.03}} \left[ \frac{E - E_e(z)}{E - 0.53 E_e(z)} \right]^{2.36} \quad (6)$$

in unit of eV-cm<sup>2</sup> following Stamnes and Rees (1983). In equation (6),  $E_e = 8.618 \times 10^{-5} T_e$  with  $T_e$  being the electron temperature in K.

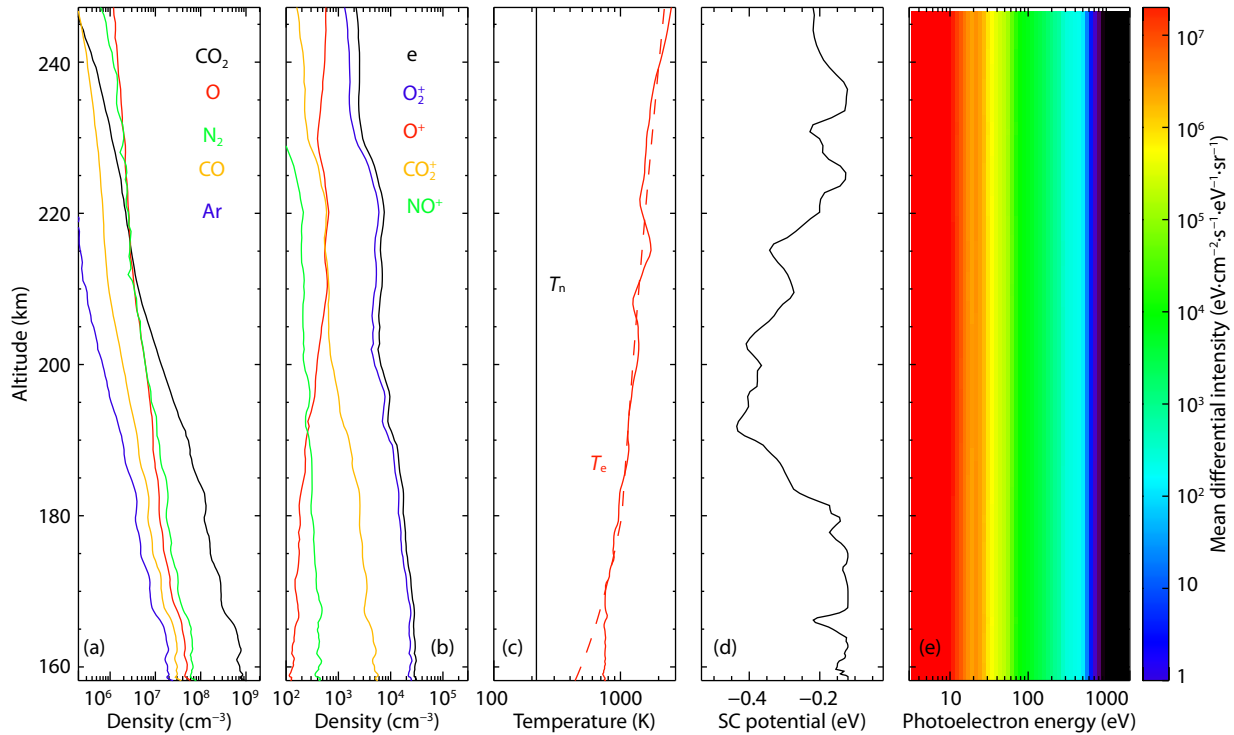
The formulism outlined above is implemented in this study assuming that all relevant parameters are isotropic (Lavvas et al., 2011). The redistribution function is parameterized similar to Lavvas et al. (2011). All necessary photon and electron impact cross sections are adapted from our previous compilation in Cui J et al. (2011).

## 3. Photoelectron Balance During MAVEN Orbit No. 2909

In this section, we verify the condition of local energy degradation for photoelectrons with the aid of the multi-instrument MAVEN data set, taking orbit No. 2909 on 28 March 2016 as an example. The periapsis altitude of this orbit is 158 km at a SZA of 67°. The 10.7 cm solar radio index at Earth during the period is 87 in solar flux unit (10<sup>-22</sup>W-m<sup>-2</sup>), indicating low solar activity conditions. The sampled regions are within the northern hemisphere of the planet characterized by draped magnetic field lines (e.g. Brain et al., 2006).

We consider a background neutral atmosphere composed of CO<sub>2</sub>, O, N<sub>2</sub>, and CO, with their densities extracted from the MAVEN Neutral Gas and Ion Mass Spectrometer (NGIMS) measurements in the closed neutral source mode (Mahaffy et al., 2015). The thermal electron density is assumed to be equal to the total ion density based on the NGIMS measurements in the open source ion mode (Benna et al., 2015), and the electron temperature is based on the characteristics of the current voltage relation obtained by the MAVEN Langmuir Probe and Waves (LPW) (Ergun et al., 2015). The incident solar EUV/X-ray flux is adapted from the solar spectral model at Mars constructed from the Flare Irradiance Spectral Model-Mars and calibrated with the MAVEN Ultraviolet Monitor band irradiance data (Thiemann et al., 2017). Finally, the photoelectron energy distribution is available from the mean differential intensities of energetic electrons measured by the MAVEN Solar Wind Electron Analyzer (SWEA) (Mitchell et al., 2016). Each SWEA spectrum is corrected for spacecraft charging using the potential derived from MAVEN Suprathermal and Thermal Ion Composition measurements (McFadden et al., 2015).

An example of our calculations of the photoelectron production and loss rates is presented in Figures 1 and 2. In Figure 1, we show from left to right the variations of neutral densities (panel a), ion and electron densities (panel b), neutral and electron temperatures (panel c), spacecraft charging (panel d), and energetic electron distribution (panel e) along the MAVEN trajectory. Only the inbound portion of the orbit is included to reduce possible wall



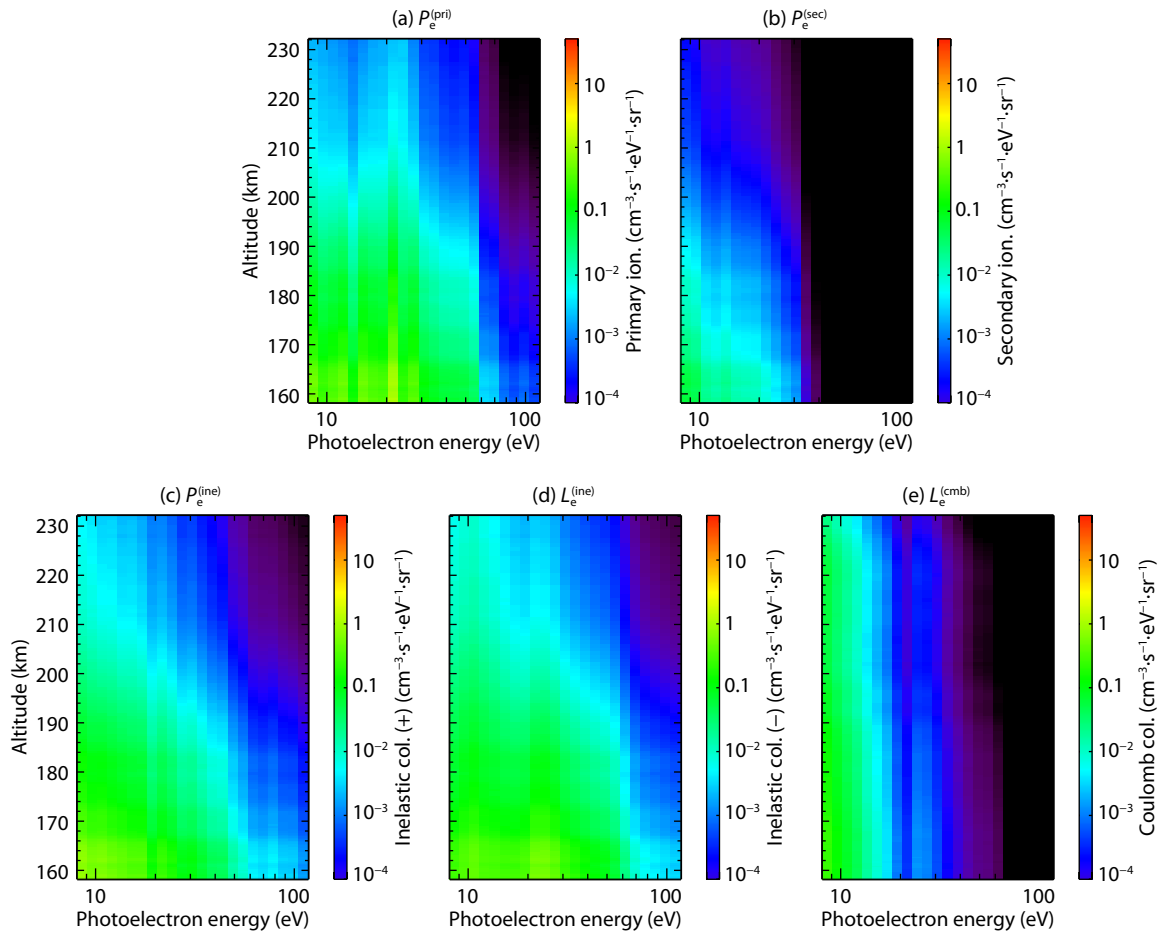
**Figure 1.** The variations of neutral densities (a), ion and electron densities (b), neutral and electron temperatures (c), spacecraft charging (d), and photoelectron energy distribution (e) along the inbound portion of MAVEN orbit No. 2909 on 28 March 2016, combining the multi-instrument data set acquired by several relevant MAVEN instruments (see text for details). The electron density is assumed to be identical to the total ion density. The neutral temperature profile is estimated from the isothermal fitting to the Ar density data; the electron temperature is corrected for the known effect of surface resistance or capacitance on the Langmuir Probe and Waves instrument sensor. The measured and corrected electron temperature profiles are distinguished as the solid and dashed lines in red in panel c.

contamination on the NGIMS antechamber walls, a well-known effect that is known to occur on other mass spectrometers with a similar design (Cui J et al., 2009). The electron temperatures measured by the LPW are known to be overestimated due to surface resistance or capacitance on the instrument sensor (Ergun et al., 2015), as indicated by the fact that the electron temperature at low altitudes is substantially higher than the neutral temperature but thermal coupling between all species is expected there (e.g. Matta et al., 2014, Peterson et al., 2018). Here the constant neutral temperature given by the solid line in black in panel c is derived from the isothermal fitting to the atmospheric Ar density profile (Stone et al., 2018). Similar to Cui et al. (2019), we adopt an empirical method to correct for the LPW measurements of electron temperature below 180 km. For comparison, the measured and corrected electron temperature profiles are indicated by the solid and dashed lines in red in panel c. The distinctive photoelectron characteristics, such as the peak structure at 22–27 eV, the knee at 60–70 eV, and another peak structure at 500 eV, are clearly seen in the photoelectron energy spectrogram in panel e.

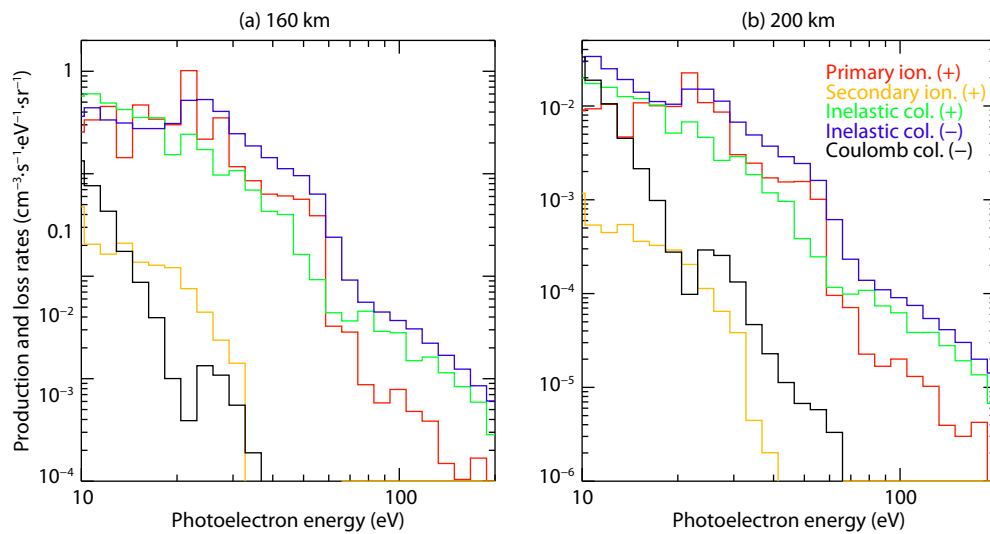
Based on the MAVEN measurements presented in Figure 1, equations (1)–(6) are used to calculate various photoelectron production and loss rates, which are displayed in Figure 2 including production via primary (panel a) and secondary (panel b) ionization, production (panel c) and loss (panel d) via inelastic collisions with atmospheric neutrals, and loss via Coulomb collisions with ionospheric thermal electrons (panel e). For a closer look, we compare

in Figure 3 the energy distributions of these rates at two representative altitudes: 160 km near the periaapsis (panel a) and 200 km (panel b). Both panels suggest that the contribution of secondary ionization is of minor importance, consistent with the relatively small ionization efficiency reported by Cui J et al. (2018) also based on the MAVEN data. For the two photoelectron loss processes, the relative importance of Coulomb collisions clearly increases with decreasing photoelectron energy. Meanwhile, it is more prominent at higher altitudes, likely connected to an increasing thermal electron to neutral density ratio which is  $3.7 \times 10^{-5}$  at 160 km and  $3.3 \times 10^{-4}$  at 200 km, respectively. According to Figures 2 and 3, the photoelectron balance is controlled mainly by primary production, as well as by production and loss via inelastic collisions with atmospheric neutrals, over the energy range considered here (see below).

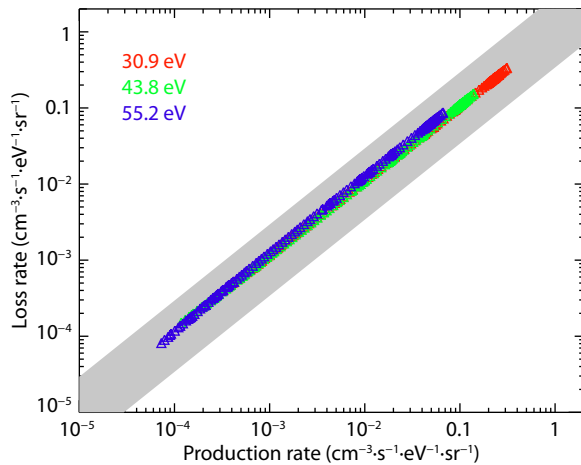
The calculated photoelectron production and loss rates are subject to several sources of uncertainty. The typical uncertainty is 10% in solar flux according to Thiemann et al. (2017) and 10% in photoionization cross section according to Heays et al. (2017), implying a combined uncertainty of 14% in primary ionization where the small uncertainty of less than 1% in neutral density is ignored. The uncertainty in differential photoelectron intensity is about 10% (Mitchell et al., 2016) and the uncertainty in electron impact cross section is about 5% (Itikawa, 2002). These lead to a combined uncertainty of 11% in secondary ionization. Following the same line of reasoning, we estimate the uncertainties in photo-



**Figure 2.** The full altitude variations of different photoelectron production and loss rates along the inbound portion of MAVEN orbit No. 2909, including production via primary (a) and secondary (b) ionization, production (c) and loss (d) via inelastic collisions with atmospheric neutrals, as well as loss via Coulomb collisions with ionospheric thermal electrons (e).



**Figure 3.** The energy distributions of the photoelectron production (indicated by “+”) and loss (indicated by “-”) rates at two representative altitudes of 160 km near the periapsis (a) and 200 km (b), both over the inbound portion of MAVEN orbit No. 2909. Different colors represent different processes including primary and secondary ionization, production and loss via inelastic collisions with atmospheric neutrals, as well as Coulomb collisions with ionospheric thermal electrons.



**Figure 4.** The relation between photoelectron production and loss rates over the inbound portion of MAVEN orbit No. 2909. Different colors stand for different photoelectron energies, as indicated in the figure legend. The region encompassing the combined measurement uncertainty is shadowed for reference.

electron production and in loss via inelastic collisions both to be 11% as well. The uncertainty in loss via Coulomb collisions is primarily determined by the uncertainty in thermal electron density, which is 25% according to Benna et al. (2015). The uncertainty in electron temperature, though as large as 20% according to Ergun et al. (2015), is not important here due to the insensitivity of the stopping function to this parameter (see equation (6)). From all the above, we conclude that the uncertainties in total production and loss to be 21% and 27%, respectively. However, we caution that a small uncertainty in spacecraft charging would lead to large uncertainties in the calculated production and loss rates at energies where an inherent rapid change in differential photoelectron intensity is expected, i.e. near the peaks at 22–27 eV and near the knee at 60–70 eV. To avoid such an undesired situation, we evaluate below the condition of local energy degradation over a restricted energy range — from 30 to 55 eV only.

Figure 2 reveals that, at 160 km, the total photoelectron production rate, which is the sum of  $P_e^{\text{pri}}$ ,  $P_e^{\text{sec}}$ , and  $P_e^{\text{ine}}$ , is generally consistent with the total loss rate, which is the sum of  $L_e^{\text{ine}}$  and  $L_e^{\text{cmb}}$ , with a median difference of 17% for any given energy between 30–55 eV. At 200 km, the agreement between production and loss is improved, with a median difference of just 13%. Both are well below the combined measurement uncertainties quoted above, implying that the condition of local energy degradation is satisfied. A more detailed comparison between the photoelectron production and loss rates is presented in Figure 4 over the altitude range of 160 to 250 km for several arbitrarily chosen energies (of 30.9 eV, 43.8 eV, and 55.2 eV, respectively) where the region encompassing the combined measurement uncertainty is shadowed for reference. The figure clearly indicates that local energy degradation is a very good approximation in the dayside Martian ionosphere at altitudes considered here.

Interestingly, the analysis presented here reveals no evidence for photoelectron transport over the considered energy range, implying that either transport occurs within the more tenuous part of

the Martian upper atmosphere (e.g. Cui J et al., 2011) or alternatively a large potential drop is present in the distant regions to reflect photoelectrons escaping back towards the planet (e.g. Kitamura et al., 2015). A thorough investigation of such an issue clearly relies on the evaluation of local energy degradation over the full energy range measured by the SWEA, which is left for follow-up studies.

#### 4. Concluding Remarks

Photoelectrons are created by solar EUV/X-ray ionization and serve as an important link between solar radiative energy and atmospheric thermal energy (Fox et al., 2008, and references therein). Observations have been made of photoelectrons near many Solar System objects such as Earth, Mars, Venus, and Titan (Coates et al., 2011). At altitudes where transport is negligible, the balance of photoelectrons is controlled by local energy degradation, a condition for which production is strictly balanced by loss (e.g. Lavvas et al., 2011).

In this study, we attempt to verify the condition of local energy degradation with the aid of the combined MAVEN measurements that provide a complete set of the physical parameters required for calculating various photoelectron production and loss rates. Taking MAVEN orbit No. 2909 as an example, our calculations indicate that photoelectron production occurs mainly via primary ionization as well as degradation from higher energy states via inelastic collisions with ambient neutrals, and photoelectron loss occurs mainly via degradation towards lower energy states also via inelastic collisions, but the effect of Coulomb collisions with ambient thermal electrons becomes important for photoelectrons less energetic than 10 eV. The condition of local energy degradation is evaluated over the energy range of 30–55 eV, chosen to reduce the influence of the uncertainty in spacecraft charging. Our analysis reveals that such a condition is very well satisfied for dayside photoelectrons up to at least 230 km. It will be interesting to examine in follow-up studies the questions of at what altitudes and what energies the condition of local energy degradation is broken and transport becomes important.

#### Acknowledgments

J Cui acknowledges supports from the National Natural Science Foundation of China (NNSFC) through grants 41525015 and 41774186. The data used in this study are publicly available at the MAVEN Science Data Center (<https://lasp.colorado.edu/maven/sdc/public/>).

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